

ASAC 1101 : FUNDAMENTALS OF SOIL SCIENCE (2+1)

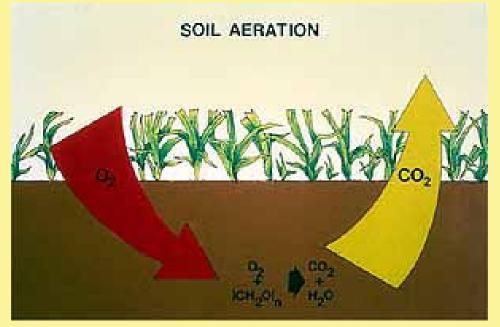
Level : B.Sc (Ag), I semester

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> **Topic** Soil Air

Soil aeration

□Soil aeration facilitates exchange of gases between soil and atmosphere

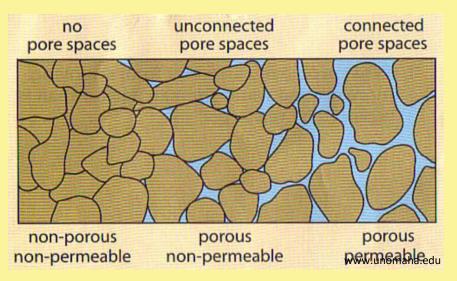


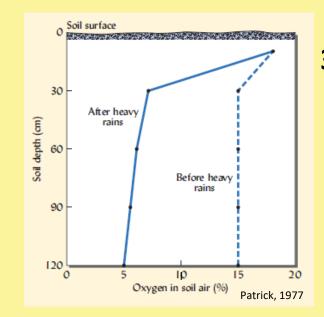
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Soil aeration

Oxygen availability in soil is determined by

1. Macro porosity





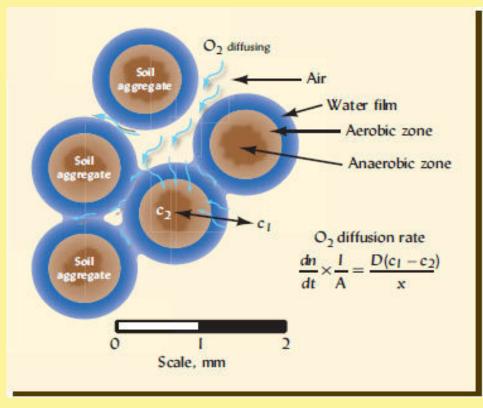
3. Soil water content

2. O₂ consumption by organisms

Soil aeration

• Oxygen availability in soil is determined by

3. Soil water content

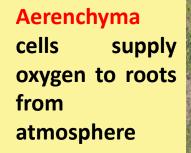


Excess moisture



Ray R. Weil

Adaption for poor aeration in waterlogging conditions

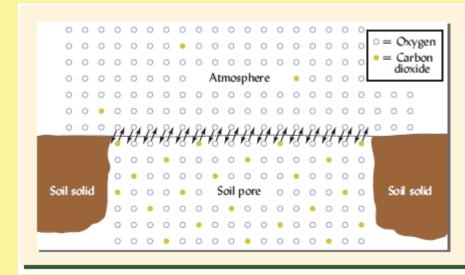




Hydrophytes are adapted to grow in waterlogged conditions
 Have above water roots to supply oxygen

Process of gaseous exchange

- Gaseous exchange occurs by mass flow and diffusion
- □ Mass flow caused by fluctuations in soil moisture content that force air in and out
- In diffusion, movement of a gas and its direction is governed by partial pressure of the gas
- **Partial pressure is the pressure exerted by an individual gas in a mixture of gases**
- **D** Partial pressure gradient causes movement of gases from high to low concentration



Calculation of partial pressure

What is the partial pressure of oxygen?

Air pressure = 100 kPa (~1 atm)
 Concentration of oxygen in air = 21%= (0.21 L/L) by vol
 Hence, partial pressure of oxygen = 21 kPa

Characterizing soil aeration

1. Gaseous composition of soil air

Soil air composed of various gases

\BoxSoil air has same amount of N₂ as atmosphere (78%)

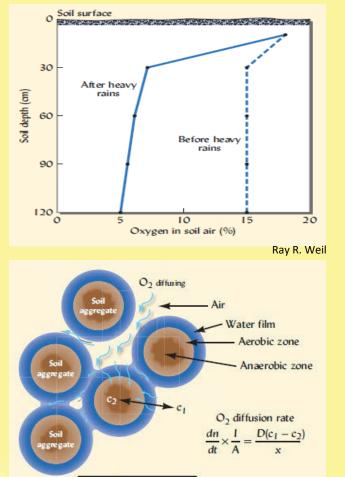
 $\Box O_2$ is less(<u>~20%</u>) in soil than atmosphere and decreases further with depth

 $\Box CO_2$ is more in soil (0.35%) (10 times more concentrated than atmosphere)

Lack of oxygen causes anaerobic conditions

 \Box Under such conditions, concentration of CH₄, H₂S and C₂H₄ increases

Characterizing soil aeration



1. Gaseous composition of soil air 2. Air filled porosity More the air filled porosity, better the aeration oxygen diffuses 10,000 slower through water filled pores than in air filled pores

Characterizing soil aeration

- 1. Gaseous composition of soil air
- 2. Air filled porosity
- 3. Oxidation-reduction (redox) potential
 - Redox potential tendency or potential for electrons to be transferred from one substance to another (E_h)

QReference for E_h : Hydrogen, $E_h = 1$

Redox potential determines which compound is present in the soil at a particular time, its concentration and oxidation state

Redox potential

Living organisms release electrons while oxidizing carbon for energy

O₂ acts as an acceptor of these electrons (oxidizing agent)

 \Box Once all O₂ is depleted by this process, anaerobic zone created

$$\stackrel{(2+)}{\underset{\text{Fe(II)}}{\overset{(2+)}{\xrightarrow{}}}} 2 \stackrel{(3+)}{\underset{\text{Fe(III)}}{\overset{(2+)}{\xrightarrow{}}}} 2 \stackrel{(3+)}{\underset{\text{Fe(III)}}{\overset{(3+)}{\xrightarrow{}}}} 2 \stackrel{(3+)}{\underset{(3+)}{\overset{(3+)}{\xrightarrow{}}}} 2 \stackrel{(3+)}{\underset{(3+)}{\overset{(3$$

- As reaction proceeds to the right, each Fe(II) loses an electron (e-) to become Fe(III) and forms H+ ions by hydrolyzing H₂O. These H+ ions lower the pH.
- When the reaction proceeds to the left, FeOOH acts as an electron acceptor and the pH rises as H+ ions are consumed.
- The tendency or potential for electrons to be transferred from one substance to another in such reactions is termed the redox potential (Eh)
- Eh measured by platinum electrode attached to a millivolt meter

Redox potential: influence of O2

Living organisms release electrons while oxidizing carbon for energy
 O₂ acts as an acceptor of these electrons (oxidizing agent)
 O₂ is in turn reduced
 Once all O₂ is depleted by this process, anaerobic zone created

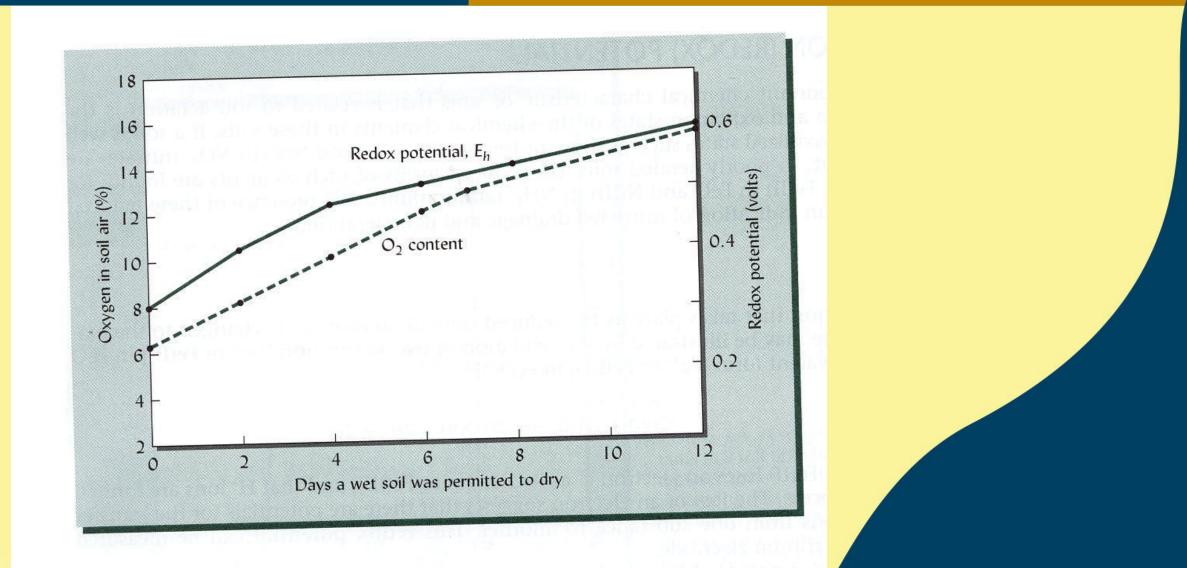
overall effect of oxidation and reduction

 $2\text{FeO} + 2\text{H}_2\text{O} \rightleftharpoons 2\text{FeOOH} + 2\text{H}^+ + 2\text{e}^ 4/2\text{O}_2 + 2\text{H}^+ + 2\text{e}^- \rightleftharpoons \text{H}_2\text{O}$

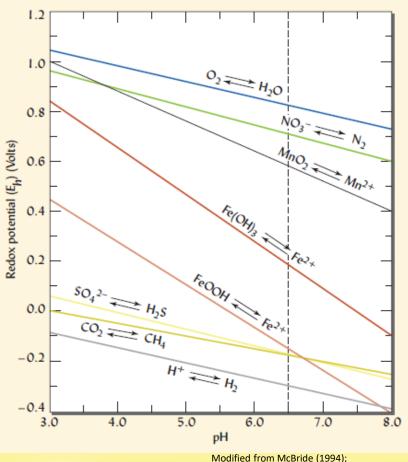
 $2\text{FeO} + \frac{1}{2}\text{O}_2 + \text{H}_2\text{O} \implies 2\text{FeOOH}$

Redox potential: influence of O2

- □In a well-aerated soil with plenty of gaseous O₂, the Eh is in the range of 0.4–0.7 volt (V).
- □As aeration is reduced and gaseous O₂ is depleted, the Eh declines to about 0.3–0.35 V.
- □If organic matter—rich soils are flooded under warm conditions, Eh values as low as -0.3 V can be found.



Relative order of reduction in redox reactions



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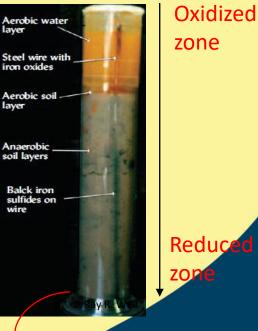
Element	Oxidized form	Charge on oxidized element	Reduced form	Charge on reduced element	change of form occurs, V
Oxygen	O ₂	0	H ₂ O	-2	0.38 to 0.32
Nitrogen	NO_3^-	+5	N ₂	0	0.28 to 0.22
Manganese	Mn ⁴⁺	+4	Mn ²⁺	+2	0.22 to 0.18
Iron	Fe ³⁺	+3	Fe ²⁺	+2	0.11 to 0.08
Sulfur	SO4 ²⁻	+6	H ₂ S	-2	-0.14 to -0.17
Carbon	CO ₂	+4	CH ₄	-4	-0.20 to -0.28

E_h values from Patrick and Jugsujinda (1992).

- **Reactions occur from top to bottom order**
- Anaerobic organisms must use electron acceptor other than O₂

E. at which

Example : iron is used here which is reduced to iron sulfide (black color)



Factors affecting soil aeration

1. Drainage of excess water

Drainage of gravitational water out of the profile and concomitant diffusion of air into the soil takes place most readily in macropores
 Dependent on soil texture, bulk density, organic matter content etc

2. Rates of respiration in soil

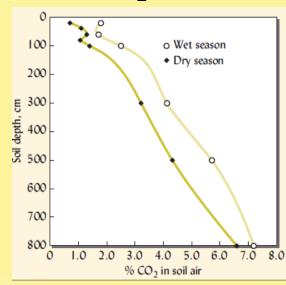
Depends on organic matter content

More the organic matter, more the microbes use them and release gases into the soil

Factors affecting soil aeration

3. Soil profile

$\Box O_2$ tends to decrease with depth $\Box CO_2$ increases with depth



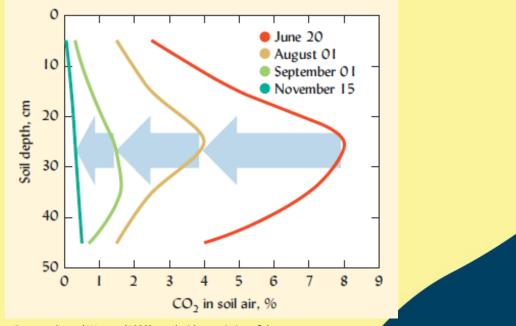


4. Soil heterogeneity Long term tillage reduces aeration Clay soil has less aeration Interped zones have more aeration than within peds

5. Seasonal differences

6. Effects of vegetation

Vegetation transpires, lowers the ground water table, improves aeration



Buyanovsky and Wagner (1983); used with permission of the Soil Science Society of America

Ecological effects of soil aeration

- Effects on organic residue degradation
 More O₂, more the decay of organic matter
 Entire soil property depends on soil aeration
- **2.** Oxidation reduction of elements

OXIDIZED AND REDUCED FORMS OF SEVERAL IMPORTANT ELEMENTS				
Element	Normal form in well-oxidized soils	Reduced form found in waterlogged soils		
Carbon	CO ₂ , C ₆ H ₁₂ O ₆	CH ₄ , C ₂ H ₄ , CH ₃ CH ₂ OH		
Nitrogen	NO ₃	N ₂ , NH ⁺ ₄		
Sulfur	SO ₄ ⁻	H ₂ S, S ^{2–}		
Iron	Fe ³⁺ [Fe(III) oxides]	Fe ²⁺ [Fe(II) oxides]		
Manganese	Mn ⁴⁺ [Mn(IV) oxides]	Mn ²⁺ [Mn(II) oxides]		

Ecological effects of soil aeration

Oxidized zones have red colors while reduced zones have grey, blue colors Anaerobic conditions causes emission of green house gases CH_4 and N_2O

3. Effects on activities of higher plants

Poor aeration affects shoot growth more than root growth
 Different plants have different degree of tolerance to waterlogging
 Low aeration also affects the nutrient uptake by plants

Fick's law

The diffusion process can be described by Fick's law

qd = -D dc/dx

qd =diffusive flux (mass diffusing across a unit area per unit per unit time)

- **D** = diffusion coefficient (area per time)
- c = concentration (mass of diffusing substance per volume)
- x = distance, and
- dc/dx = the concentration gradient

If partial pressure p is used instead of concentration of the

diffusing component, we get qd = -(D/ß)(dp/dx)

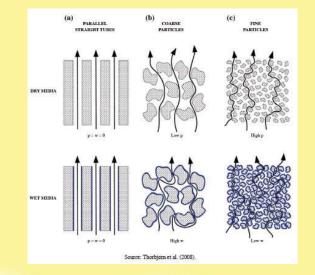
ß = ratio of the partial pressure to the concentration

Fick's law

Considering first the diffusive path in the air phase,

➤ the diffusion coefficient in the soil D_s must be smaller than that in bulk air D₀ owing to the limited fraction of the total volume occupied by continuous air-filled pores and also to the tortuous nature of these pores

> Ds to be some function of the air-filled porosity, fa .



Fick's law

Different workers have over the years found different relations between Ds and fa for various soils. For instance, Buckingham (1904) reported the following nonlinear relation: Ds/ D₀ = Kfa²

> Penman (1940) found a linear relation : Ds/ D₀ = 0.66 fa

0.66= tortuosity coefficient, suggesting that the <u>apparent path is</u> <u>about 2/3 the length of the real average path of diffusion in the soil</u>

Reference:

The Nature and Properties of Soils by Nyle C. Brady and Ray R. Weil